

# Absorption-Based Size-Specific Primary Productivity Algorithm for the River-Influenced Northern Gulf of Mexico

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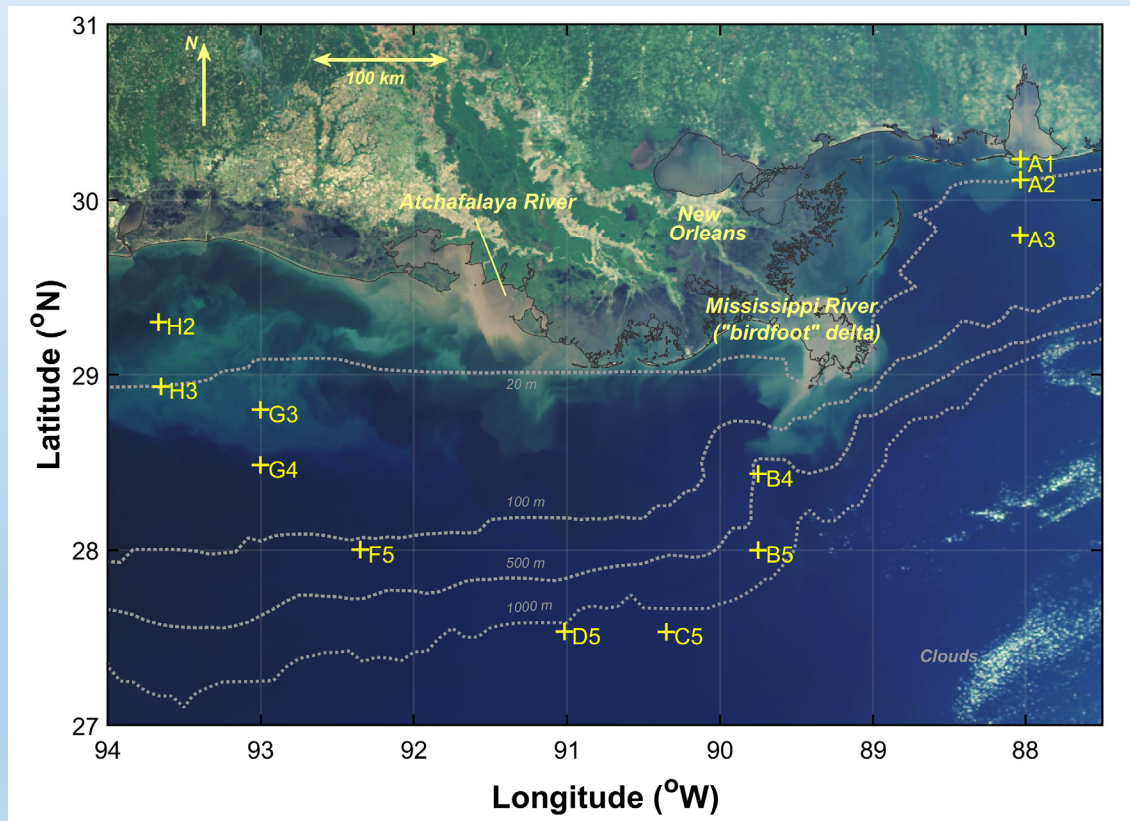
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- Introduction
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# Introduction – Study Area and Data

- Large river system associated with the Mississippi and Atchafalaya rivers and one of the largest signals for carbon cycling in the North American continent



- Global PP models may not be able to realistically simulate regional conditions
- Here, we examine a regional absorption-based and size-specific productivity algorithm that can be implemented using satellite-retrievable observations
- Cruise conducted in April 2009
- Stations included coastal, mid-shelf and slope waters
- Dataset included:
  - P-E measurements
  - Quantitative filterpad
  - Hyperspectral irradiance profiles

# Algorithm – Absorption-based and size-specific

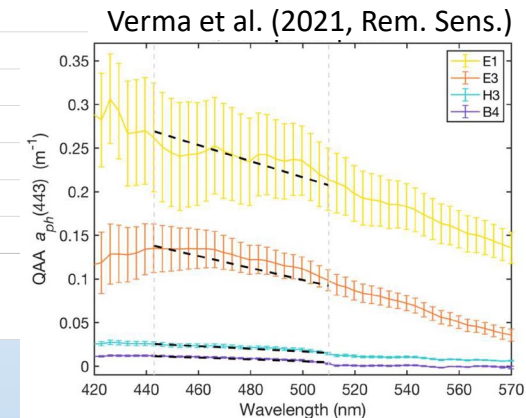
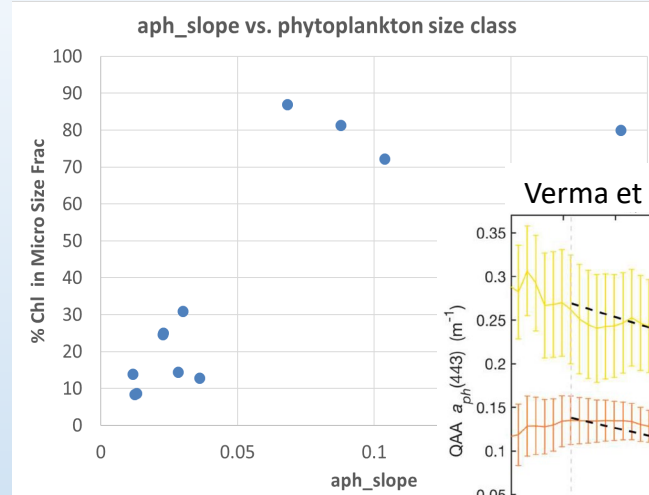
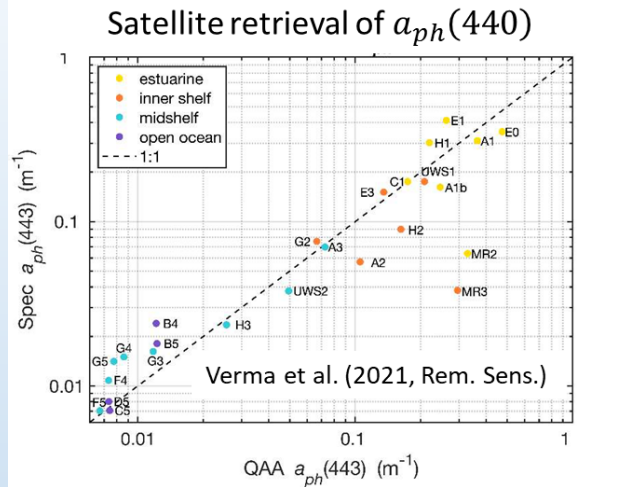
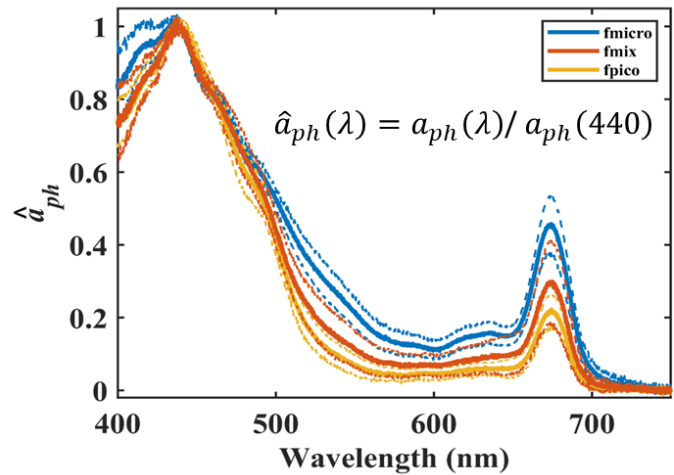
- $$P(z, t) = \bar{a}_{ph} P_{max}^{a_{ph}} \left[ 1 - \exp \left( - \frac{a_{ph}(440) \phi_{max}^C PUR(z, t)}{\bar{a}_{ph} P_{max}^{a_{ph}}} \right) \right] \text{ (mol C m}^{-3} \text{ h}^{-1}\text{)}$$
  - $$PUR(z, t) = \int_{PAR} \hat{a}_{ph}(\lambda) E_d(z, \lambda, t) d\lambda \text{ where } \hat{a}_{ph}(\lambda) = a_{ph}(\lambda) / a_{ph}(440)$$
  - $$\bar{a}_{ph} = \frac{\int_{400}^{700} a_{ph}(\lambda) E_d(z, \lambda, t) d\lambda}{\int_{400}^{700} E_d(z, \lambda, t) d\lambda} \text{ where } a_{ph}(\lambda) = \hat{a}_{ph}(\lambda) a_{ph}(440) \text{ (m}^{-1}\text{)}$$
  - $$\phi_{max}^C = 0.345 a_{ph\_slope}^2 + 0.195 a_{ph\_slope} + 0.017 \text{ (mol C (mol quanta)}^{-1}\text{)}$$

where  $a_{ph\_slope} = \frac{[a_{ph}(443) - a_{ph}(510)]}{510 - 443} \text{ (m}^{-1} \text{ nm}^{-1}\text{)}$
  - $$P_{max}^{a_{ph}} = Chl \cdot P_{max}^B / \bar{a}_{ph} = 2.00 \times 10^{-5} e^{0.306(T)} \text{ (mol C m}^{-2} \text{ h}^{-1}\text{)}$$
- $$IP = \int_0^{DL} \int_0^D P(z, t) dz dt \text{ (mol C m}^{-2} \text{ d}^{-1}\text{)}$$

where  $DL$  is daylength and  $D$  is water column depth

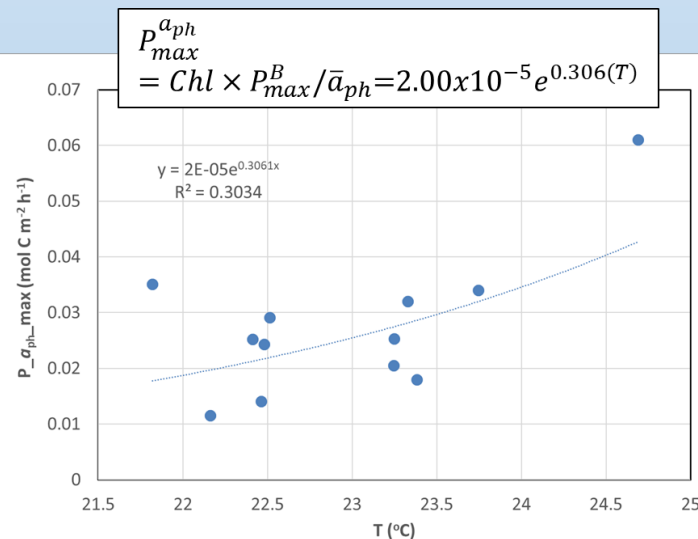
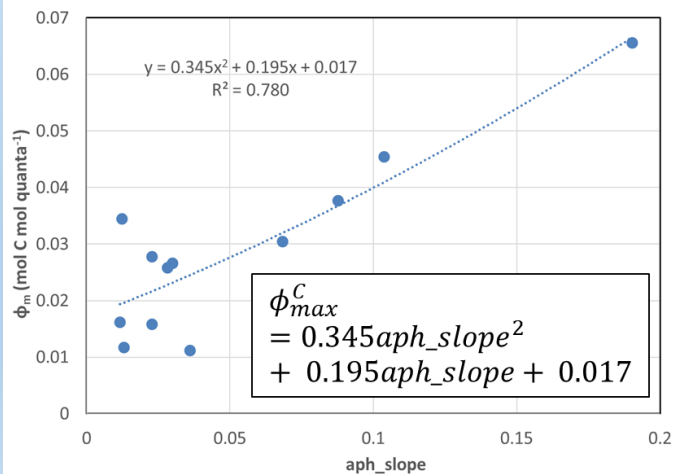
# Absorption shape vectors and $a_{ph}(440)$

# Model Inputs

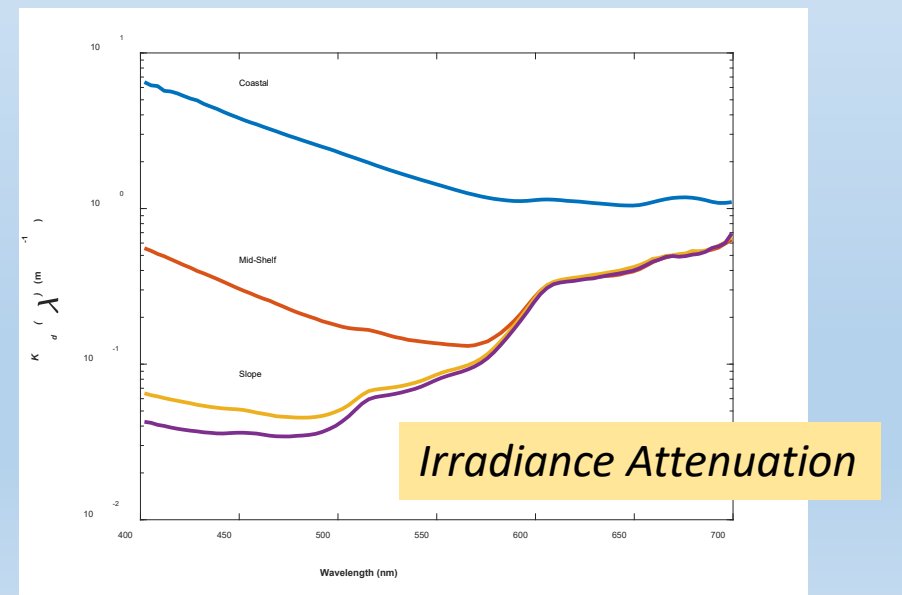


$$\bar{a}_{ph} = \frac{\int_{400}^{700} a_{ph}(\lambda) E_d(z, \lambda, t) d\lambda}{\int_{400}^{700} E_d(z, \lambda, t) d\lambda} \text{ where } a_{ph}(\lambda) = \hat{a}_{ph}(\lambda) a_{ph}(440) (m^{-1})$$

## Absorption slope

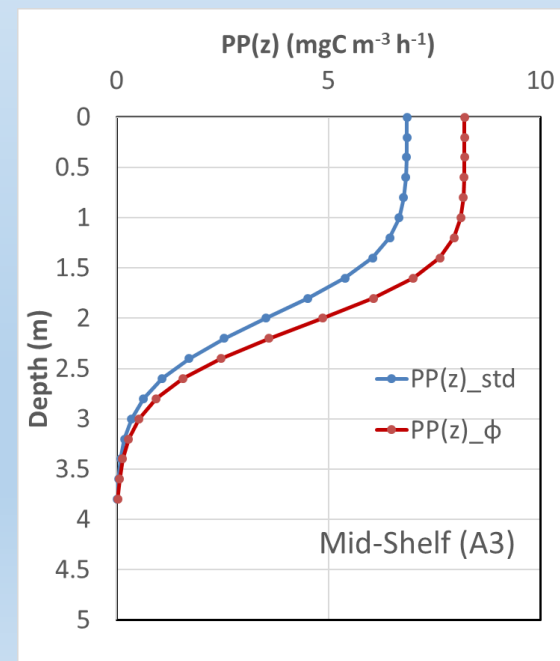
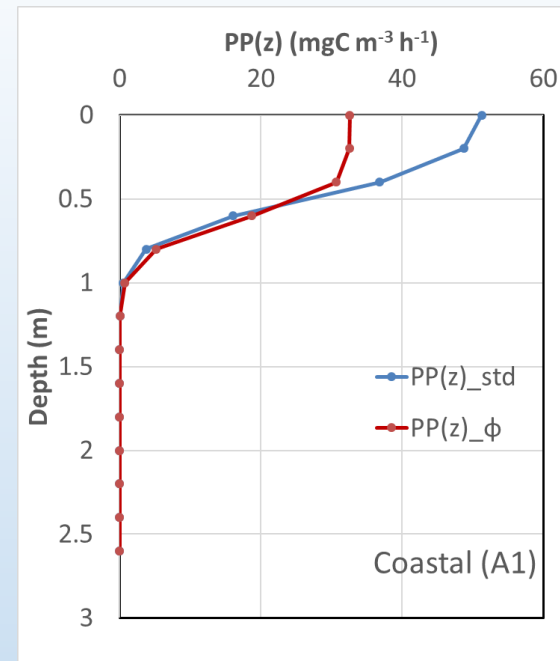


## Photophysiological parameters

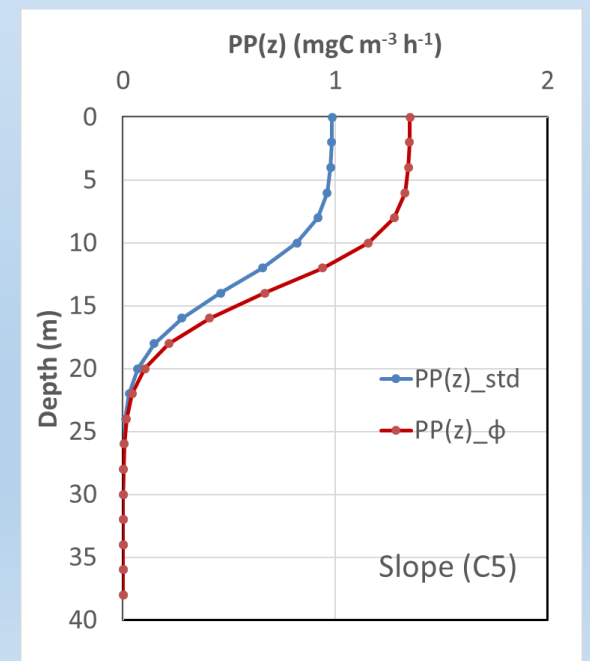


# Results and Conclusions

- Preliminary demonstration of model shows it is able to represent order of magnitude variations in productivity from nearshore to slope waters
- Largest deviations were in near surface waters, where sensitivity to maximum photosynthetic rate estimates were highest
- Positive bias in model for slope and mid-shelf and negative for coastal



**PP(z)\_ $\phi$  = model**  
**PP(z)\_std = measured**



# Knowledge gaps and priorities

- Knowledge gaps
  - Large source of uncertainty in absorption-based models in the maximum photosynthetic rate parameter
  - Need for additional in situ measurements to better characterize PP parameters
- Priorities
  - 1 year
    - Expand analysis to include datasets from other seasons and conditions to provide more comprehensive representation of photophysiological parameters
    - Evaluate performance of model using satellite-retrieved parameters ( $a_{ph}(440)$ ,  $aph\_slope$ ,  $K_d$ , and T)
  - 5 year
    - Apply algorithm to emerging hyperspectral sensors (PACE, GLIMR, SBG)
    - Compare model performance to other formulations including global algorithms
  - 10 year
    - Explore performance and suitability of algorithm in other regions, and potential “nesting” of models